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Discordant Silurian paleolatitudes for central Newfoundland: New paleomagnetic evidence from the Springdale Group

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ABSTRACT

Ancient remanences are retained by the Early Silurian (429 ± 6/–5 Ma) mafic volcanics of the Springdale Caldera (five sites) and the overlying red beds (seven sites). Dual polarity magnetizations are obtained by thermal demagnetization of samples from the red beds, whereas single polarity directions are observed in the volcanics. High unblocking temperatures indicate hematite as the remanence carrier in both the volcanics and sediments. These high-temperature, characteristic remanences are easily isolated and pass both the tilt and conglomerate tests, they are likely to be of primary Silurian age. Characteristic declinations are predominantly northerly and northeasterly, and indicate significant structural rotations on a local scale. When the results of the red beds and the volcanics are combined they show characteristic inclinations that are shallower than those of the correlative Botwood Group (ca. 36° vs. 43°) but not nearly as shallow as those reported from the King George IV Lake area (0.5°). Mean inclinations obtained from the Springdale red beds are, however, significantly shallower than those of the Springdale volcanics. The same difference can be seen in other previous Silurian paleomagnetic studies of central Newfoundland. We infer that an inclination error affects the red bed magnetizations of the Springdale Group, Botwood Group (Wigwam Formation) and rocks of the King George IV Lake area. Therefore, the results from Silurian red beds should not be used to determine paleolatitudes for central Newfoundland. The mean paleolatitude of the Springdale Group volcanics is 30°. The mean paleolatitudes for both the Springdale volcanics and Botwood volcanics (Lawrenceton Formation) are indistinguishable within paleomagnetic error limits from the predicted paleolatitude of Newfoundland on the northeast-trending North American margin. Thus, no detectable post-Silurian displacement is shown by the volcanics of the Springdale Group with respect to cratonic North America.

1. Introduction

The northern Appalachians of Newfoundland may be divided into three main regions, the Humber Zone (Laurentian Margin), the Avalon Zone, and the intervening Central Mobile Belt (CMB). These zones represent the margins of the Ordovician Iapetus Ocean and its remnants. The CMB may be further subdivided into the Dunaige Zone, consisting of the Exploits and Notre Dame Subzones, and the Gander Zone [1], but for the purpose of this study we will use the term CMB (Fig. 1).

Paleomagnetic investigations of the Silurian overlap assemblages in central Newfoundland have yielded shallow and intermediate inclinations. Shallow inclinations are exemplified by the

results from the King George IV Lake area [2] and by the results from the Wigwam Formation of the Botwood Group [3, 4]. Intermediate inclinations have been observed from the Lawrenceton Formation of the Botwood Group [3, 5]. Paleomagnetic studies of these Silurian units can be useful for the documentation of the timing and nature of displacement associated with the closure of Iapetus, but the inclination discrepancies cast doubt on the reliability of these results and they must be resolved before conclusions can be drawn.

Paleolatitudes obtained from Silurian units indicate either significant (~16°) southward (dextral) movement with respect to Laurentia (King George IV Lake area and Wigwam Formation) or no detectable displacement (Lawrenceton Forma-

tion) of these cover sequences. The significant dextral offset suggested by the shallow directions is of the order of thousands of kilometers and is thus amenable to paleomagnetic documentation. It is important to note that steeper inclinations are predominantly associated with volcanics (Lawrenceton Formation) [3,5], whereas the shallower inclinations are often associated with red beds (King George IV Lake area and Wigwam Formation) [2,3,4]. These inclination discrepan-

cies would suggest largely different paleolatitudes for coeval Silurian rocks, which is implausible.

In this study we attempt to clarify and resolve this issue, by including both the mafic volcanic rocks of the Silurian Springdale Caldera and the overlying Springdale Group red beds in our paleomagnetic investigation. Whereas the structural corrections can be well determined for the red beds and somewhat less so for the volcanics, the timing and nature of remanence acquisition may

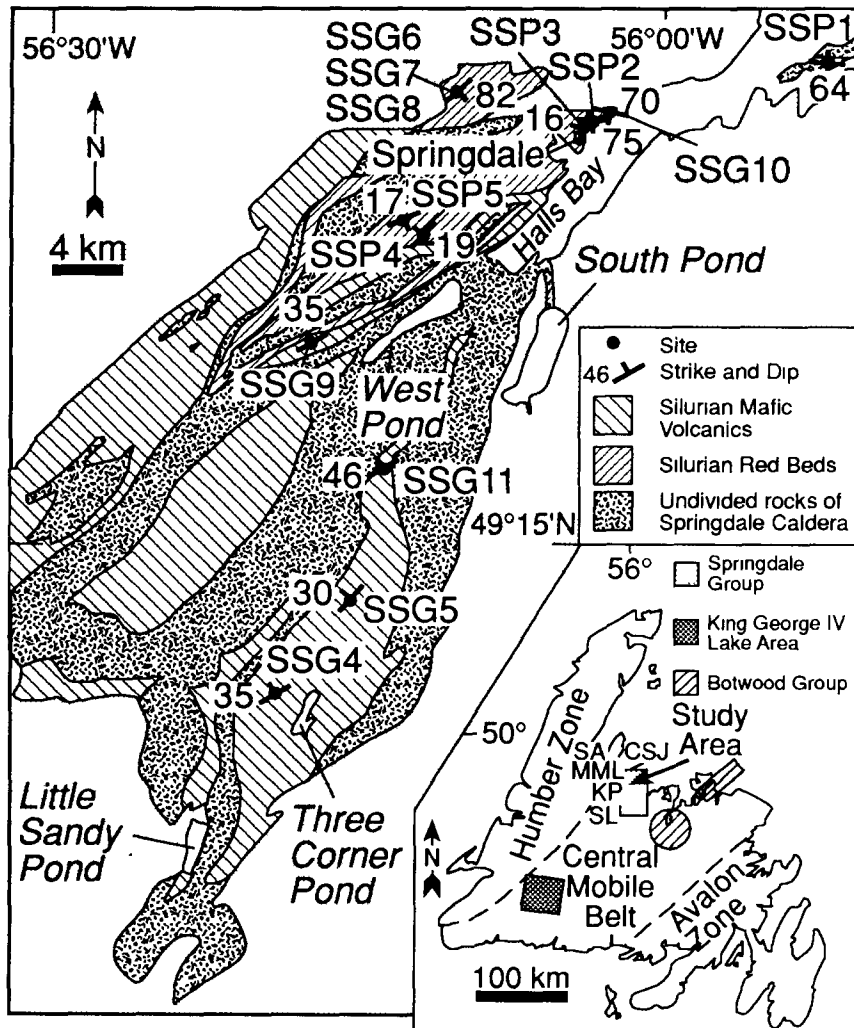


Fig. 1. Simplified geologic map of the Springdale Caldera [11], with paleomagnetic site locations. Inset. Location map of the Springdale Group study area within the Central Mobile Belt of Newfoundland. *CSJ* = Cape Saint John Group, *MML* = Mic Mac Lake Group, *KP* = King's Point Complex, *SL* = Sheffield Lake Group, *SA* = Sops Arm Group. Together with the Springdale Group, these correlative Early Silurian volcanic centers form an overlap assemblage linking the Humber Zone (Laurentian margin) with the western Central Mobile Belt [10,9].

be better known for the volcanic rocks. A previous paleomagnetic study of the Springdale Group is the unpublished study of Black [1965, as cited in 2]. This included samples of both volcanics and red beds and resulted in a mean direction of $D = 17.7^\circ$, $I = -18.4^\circ$ ($k = 24.7$, $\alpha_{95} = 5.6^\circ$) with a corresponding paleomagnetic pole located at 103.7°E , 29.1°N ($\delta_p = 5.8^\circ$, $\delta_m = 3.0^\circ$). This result must be considered incomplete because it included no more than 28 samples and only AF demagnetization was performed to a maximum of 30 mT; it is therefore unlikely that the characteristic remanence, carried by hematite, was isolated. Also, the age of this magnetization is unconstrained as no field tests are reported.

The Springdale Group unconformably overlies the Ordovician Roberts Arm Group [6,7], which has recently been studied by Van der Voo et al. [8]. This submarine volcanic unit has a paleolatitude distinctly different from that of Laurentia (31° vs. 15° , respectively). The Springdale Group appears to be correlative with surrounding Silurian units: to the northwest, the King's Point Complex; to the north on the Baie Verte Peninsula, the Cape St. John Group; to the west, the Mic Mac Lake and Sheffield Lake Groups; to the west at the southern end of the Humber Peninsula, the Sops Arm Group [9]. The continuity of lower Silurian units in western Newfoundland suggests a linkage of the western Central Mobile Belt to the North American margin by this overlap assemblage [9,10]. This study was undertaken firstly to constrain the timing of the docking of CMB terranes in a test of the linkage/overlap situation suggested by the geology, and secondly to examine the disparate inclination results and corresponding paleolatitudes obtained from Silurian red beds and volcanics of the northern Appalachians.

2. Geologic setting

The Springdale epicontinental-type collapse caldera extends at least 60 km along strike and has a maximum width of 35 km (Fig. 1). This elongate outcrop pattern is parallel to the north-northeast regional fault orientation. The youngest volcanic unit, a nearly massive felsite (unit 10 [11]), is of Early Silurian age based on zircon U-Pb ages of $429 + 6 / - 5$ Ma and 425 ± 3 Ma by

Chandler et al. [10] and Coyle and Strong [11] respectively. This youngest volcanic unit is interbedded with red bed sandstone (unit 9 [11]) and is located in the northern section of the outcrop area. This uppermost unit thus yields a minimum age for the Springdale Group. The oldest volcanic unit (unit 1 [11]) within the Springdale Caldera yields a maximum age for the Springdale Group of $432.4 + 1.7 / - 1.4$ Ma based on U-Pb zircon geochronology [11]. Thus, the volcanoplutonic events of the Springdale volcanic field took place within a 5 m.y. period between a maximum of 434 and a minimum of 425 Ma [11].

The Springdale Group comprises basaltic flows, andesite flows and pyroclastic rocks, silicic ash-flow tuffs, high-silica rhyolite domes, volcanically derived debris flows and breccias, and fluvialite red sandstones and conglomerates [9]. The magmatic activity is envisaged as occurring after the closure of this section of Iapetus in an overall transpressional–extensional tectonic regime [9]. Most of the caldera volcanics are not well exposed; thus we have concentrated our efforts on the mafic volcanics and overlying red beds. The detailed petrology and geochemistry of the Springdale Caldera is presented in Coyle [12] and the sedimentology of the overlying red beds is described by Wessel [13]. To the east and west, the caldera is bounded by upfaulted basement rocks and is intruded by granitoid rocks of the Topsails Complex. To the northwest and northeast it unconformably overlies Lower to Middle Ordovician submarine volcanics. These include the Lush's Bight, Catcher's Pond and Roberts Arm Groups. The latter has yielded a U-Pb zircon age of 473 ± 2 Ma [14].

The Springdale Group shows a lack of penetrative deformation (cleavage) and the synform is folded about a gently plunging northeasterly axis. In the caldera, the beds generally dip gently, but locally dip as steeply as 50° [9]. Coyle and Strong [9] have estimated up to 20% of across-strike regional shortening to account for the present elongate map pattern of the Springdale Group.

3. Field methods

Structural attitudes (paleohorizontal) were determined from flows (pillows and flow tops) and interbedded and overlying well-bedded red sedi-

ments. Sites with varying strike and dip were obtained in order to perform a fold test. During the summers of 1989 and 1991, over 120 samples from thirteen sites within the Springdale Group were collected for paleomagnetic study. At least six samples were collected per site. Most samples were drilled in place with a portable gas-powered core drill, and oriented with a magnetic compass and clinometer. Additional samples were collected as oriented blocks and drilled in the laboratory. The low magnetic intensities of these rocks do not significantly affect the magnetic compass

readings, as demonstrated in the field by the lack of deflection of the magnetic compass needle.

4. Laboratory methods

The cores were cut into standard size paleomagnetic samples (2.25 cm × 2.5 cm) and stored within a magnetically shielded room at the University of Michigan Paleomagnetic Laboratory. Natural remanent magnetization (NRM) directions were measured on either a ScT two-axis superconducting rock magnetometer or a Schon-

TABLE 1
Springdale Group site means for the characteristic component

Site	<i>S/D</i>	<i>N</i>	<i>D/I in-situ</i>	<i>D/I tilt-corrected</i>	α_{95}	<i>k</i>	Paleomag pole Lat /Long
<i>Red beds</i>							
SSP1	274/64	6	59.5/19.4	59/-18	13.1	27.1	
SSP2	93/75	10	271.1/63.8	209/13	9.9	25.0	
SSP3	197/16	7	18.6/-19.3	24/-18	12.0	26.0	
SSP4	68/17	0					
SSP5	233/19	5	15/-24.8	25/-35	21.5	13.6	
SSG6	49/82	10	70.3/-24.8	24/-22	13.7	13.5	
SSG8	49/82	11	conglomerate				
SSG10	103/70	6	252.9/-48.7	329/-34	8.1	69.0	
<i>Volcanics</i>							
SSG4	240/35	12	352.2/-28.9	11/-59	3.0	207.7	
SSG5	220/30	10	350.2/-44.4	25/-61	6.0	63.0	
SSG7	49/82	7	86.2/-23.7	25/-37	31.4	4.7	
SSG9	240/35	12	58.2/-26.7	75/-23	4.2	109.6	
SSG11	240/46	14	351.3/-7.6	349/-42	6.4	39.6	
Means							
<i>Volcanics</i>		5	22.6/-33.0		44.4	3.9	
		5		28/-49	28.3	8.3	7°N/100°E
<i>Red beds</i>		6	50.1/-13.0		45.8	3.1	
		6		24/-26	23.3	9.2	24°N/98°E
<i>All sites</i>		11	38.0/-23.1		29.4	3.4	
		11		25/-36	17.1	8.1	17°N/99°E
Inclination only							
<i>Volcanics</i>		5	-26.3		20.8	19.1	
		5		-44	25.6	12.9	
<i>Red beds</i>		6	-8.2		58.4	2.2	
		6		-19	22.9	11.0	
<i>All sites</i>		11	-21.0		24	4.3	
		11		-31	17	7.8	

S/D = strike and dip (right-hand rule), *N* = number of thermally demagnetized samples used to calculate the site mean, *D/I* = declination and inclination (°), α_{95} = radius of cone of confidence [17,23], *k* = estimate of Fisher precision parameter [17,3]

stedt SSM-1A spinner magnetometer. Alternating field demagnetizations of the NRM were performed with either a Schonstedt GSD-1 single-axis AF demagnetizer or a Sapphire Instruments SI-4 AF single-axis demagnetizer. Thermal demagnetization was conducted using a Schonstedt TSD-1 furnace, and the temperatures presented have been recalculated based on a 1991 recalibration of the thermal demagnetizer. Alternating field demagnetization of pilot samples failed to resolve multiple components of remanence and left a significant portion of high-coercivity NRM intensity even after applied fields of 170 mT. We therefore performed thermal demagnetization in this study. Thermal demagnetization from room temperature (NRM) to approximately 700°C used between 15 and 30 steps.

Principal component analysis [15] was used to fit lines and planes (great circles) to the components identified by inspection of the orthogonal vector diagrams [16]. A minimum of at least three points were fit by the lines and planes and the maximum acceptable mean angular deviation (MAD) was chosen to be 15°. Site means and a formation mean are presented with their respective α_{95} cones of confidence and precision parameter k [17] in Table 1.

5. Results

Natural remanent magnetization (NRM) intensities range from 1 to 15 mA/m for the red bed samples and from 50 to 2000 mA/m for the volcanic samples. Northerly and steeply downward directions from some samples suggest a partial present day field (PDF) direction overprint, but the corresponding NRM directions differ from the present-day field direction, indicating that the characteristic remanence has not been completely overprinted by a viscous magnetization of recent origin. Other samples exhibit northerly and shallow upward directions that resemble the characteristic direction, as will be shown later.

5.1 Demagnetization

In general, the demagnetization trajectories of the volcanic samples are very well behaved, whereas those of red bed samples are more noisy,

which may account for the larger α_{95} confidence circles associated with the red bed sites. The demagnetization behavior of the red bed samples is illustrated by samples SSP2.6 and SSP3.7 (Fig. 2a and b). Sample SSP2.6 exhibits removal of a component (present-day field overprint in *in-situ* coordinates) at fairly low temperatures, followed by a univectorial trend to the origin with an unblocking temperature spectrum between 640 and 673°C. Sample SSP3.7 exhibits a curved trajectory at fairly low unblocking temperatures during the removal of a steep (present-day field?) overprint. As with the previous example the sample exhibits a range of unblocking temperatures between 625 and 673°C.

The demagnetization behavior of the volcanic samples is illustrated by samples SSG11.12 and SSG5.6 (Fig. 2c and d). Sample SSG11.12 exhibits a near-univectorial orthogonal vector projection and complete removal of the NRM is not attained until 673°C. This clean behavior is characteristic of many of the volcanic samples. Sample SSG5.6 exhibits linear demagnetization as well as high unblocking temperatures up to 687°C and is also illustrative of the demagnetization behavior of the volcanics.

Dual polarity characteristic magnetizations are observed in the red beds whereas only a single polarity has been found in the volcanics. The presence of a red bed site (SSP2) with reversed directions that are nearly antipodal to the mean direction suggests that paleosecular variation has been adequately averaged.

The characteristic magnetization is carried by hematite, as can be concluded from the distinctive high-temperature unblocking spectra and the failure of AF demagnetization to remove more than half of the NRM intensity. Also, IRM experiments fail to magnetically saturate either the red bed or volcanic samples, even at peak fields up to 1.4 T.

5.2 Conglomerate test

Paleomagnetic cores were collected from eleven clasts of volcanic rock in an upper unit of puddingstone conglomerate within the Springdale Group red beds. Individual clasts show similarities to, as well as differences from, the demagnetization behavior of individual samples of Spring-

dale Caldera volcanics. Similarities include high-temperature (hematite) unblocking spectra, whereas differences include trajectories that are curved to a greater degree at low temperatures as well as noisier high-temperature behavior.

As shown in the stereogram (Fig. 3) these high-temperature stable end point directions are randomly oriented, thus demonstrating that this unit has not been remagnetized since the time of deposition [18]. This conglomerate test is random

at the 95% significance level, as the calculated resultant ($R = 1.66$) is less than the expected value ($R_0 = 5.28$) [19,20]. The distribution of magnetization directions may be termed random, as the length of the resultant vector (R) is less than R_0 , where R_0 is the value of R that will be exceeded with a probability of 5% by a sample drawn from a random population. Thus, the characteristic high-temperature remanence direction passes the conglomerate test.

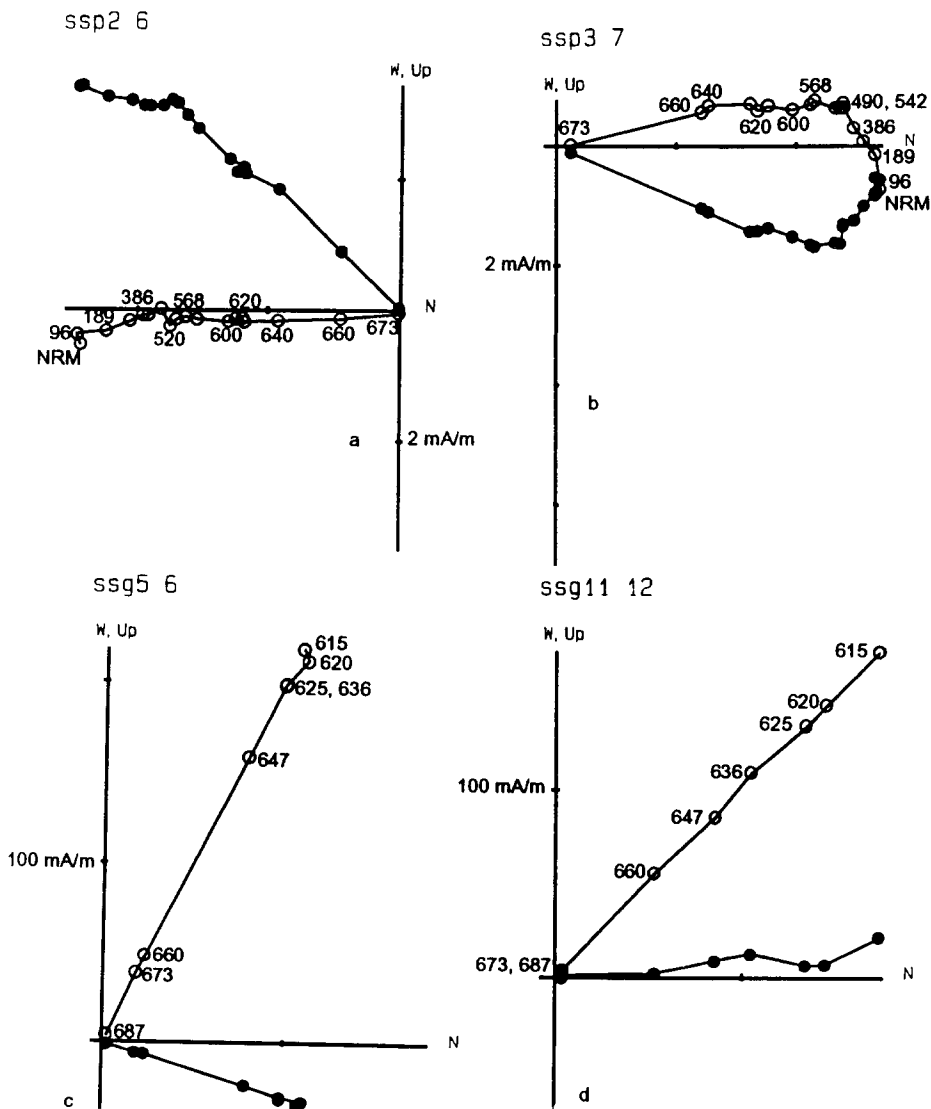


Fig 2 Tilt-corrected orthogonal vector projections of thermal demagnetization behavior for Springdale Group red bed samples SSP2 6 (a) and SSP3 7 (b) and volcanic samples SSG5 6 (c) and SSG11 12 (d) ● = projections of the horizontal component, ○ = projections of the vertical component

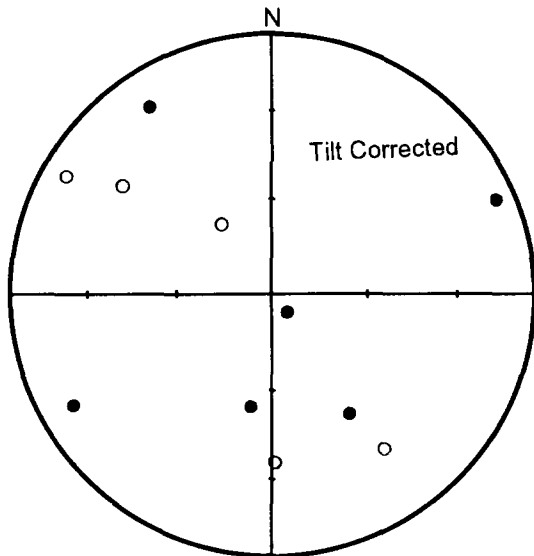


Fig 3 Tilt-corrected equal-area stereogram illustrating the positive conglomerate test for site SSG8. ● = projections on the lower hemisphere; ○ = projections on the upper hemisphere. These high-temperature stable end point directions from volcanic clasts ($N=11$) pass the conglomerate test. The calculated resultant ($R=1.66$) is less than predicted ($R_0=5.28$), where R_0 is the value that will be exceeded with a probability of 5% when sampling from a random population [19,20]

5.3 Tilt test

The site mean directions for volcanics and red beds combined cluster better after tilt correction (*in-situ*: $D/I = 38.0^\circ / -23.1^\circ$, $k = 3.4$, $\alpha_{95} = 29.4$, $n = 11$ sites; tilt-corrected: $D/I = 25^\circ / -36^\circ$, $k = 8.1$, $\alpha_{95} = 17.1$, $n = 11$). The tectonic correlation fold test of McFadden [21] yields a value ($SCOS = 3.861$) that is nearly equal to the 95% confidence value (3.865). Thus there is no reason to reject, at nearly the 95% confidence level, the hypothesis that the characteristic magnetization was acquired before folding. This suggests a pre-tilting age of magnetization (Fig. 4) and the lack of a dramatic improvement in clustering is attributable to the wide between-site scatter of declinations. We have analyzed this declination variation by comparing it with the regional variation in strike in an attempt to account for any subsequent regional warping of the Springdale Caldera, but this analysis has not yielded a coherent pattern. The calculated paleomagnetic pole

(17.3°N , 99.1°E ; $\delta_p = 11.5^\circ$, $\delta_m = 19.9^\circ$) plots close to the Siluro-Devonian reference poles [22] (Fig. 5).

The formation mean may be divided into red bed and volcanic site submeans (see Table 1). Because we believe that the declinations have been scattered by later structural disruption we have also calculated the mean inclinations without regard to declination. The method of McFadden and Reid [23] was used to calculate k , α_{95} and the mean inclinations. These mean inclinations are as follows:

For all sites: *in-situ* $-21.0 \pm 24.0^\circ$, tilt-corrected $-31 \pm 17.0^\circ$ ($n = 11$ sites)

For red bed sites: *in-situ* -8.2 ± 58.4 , tilt-corrected $-19 \pm 22.9^\circ$ ($n = 6$ sites)

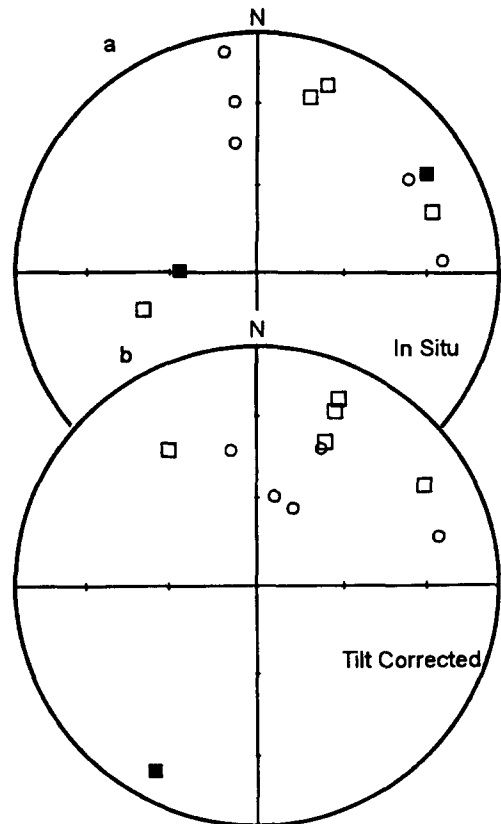


Fig. 4 Equal-area stereograms for all Springdale Group site means in *in-situ* (a) and tilt-corrected coordinates (b). *In-situ* mean $D/I = 38.0^\circ / -23.1^\circ$ ($k = 3.4$, $\alpha_{95} = 29.4$). Tilt-corrected mean $D/I = 25^\circ / -36^\circ$ ($k = 8.1$, $\alpha_{95} = 17.1$). The squares/circles correspond to red bed/volcanic sites. Solid/open symbols are projections on the lower/upper hemisphere.

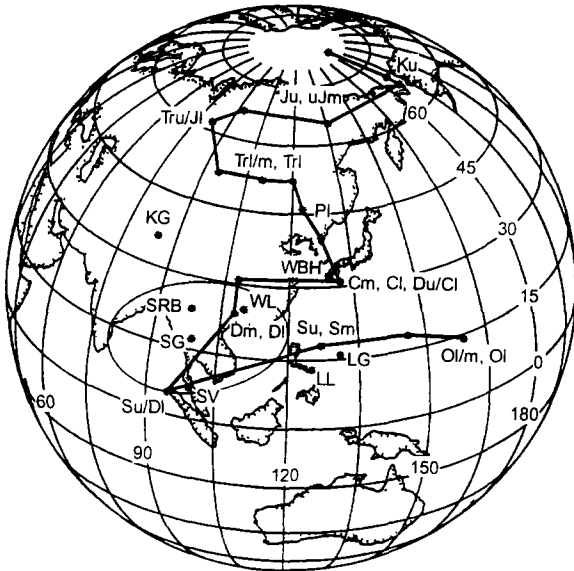


Fig 5 APWP for North America comprising mean paleopoles with $Q \geq 3$ after Van der Voo [22] LL = Lawrenceton Formation of Lapointe [3] (13°N , 125°E); LG = Lawrenceton Formation of Gales et al [5] (16°N , 131°E), KG = King George IV Lake area of Buchan and Hodych [2] (35°N , 84°E), WBH = Wigwam Formation of Buchan and Hodych [4] (32°N , 129°E), WL = Wigwam Formation of Lapointe [3] (25°N , 110°E); SRB = Springdale red beds of this study (24°N , 98°E), SV = Springdale volcanics of this study (7°N , 100°E), SG = combined Springdale Group of this study (17°N , 99°E) with confidence ellipse

For volcanic sites: *in-situ* $-26.3 \pm 20.8^{\circ}$, tilt-corrected $-44 \pm 25.6^{\circ}$ ($n = 5$ sites).

6. Discussion

The positive conglomerate and tilt tests and the presence of reversals lead us to conclude that the characteristic magnetization of the Springdale Group is a primary magnetization. In addition, the paleomagnetic pole does not closely resemble post-Devonian paleomagnetic poles from North America as might be expected if the magnetization were representative of a late remagnetization event.

While the individual site means show good agreement in inclination, the declinations vary significantly. This may be due to differential block rotations upon collapse of the caldera and the subsequent structural elongation of the caldera [9] or later (Devonian–Carboniferous) strike-slip

movements [11]. For our analysis we are concerned only with the paleolatitude of the Springdale Group, which is derived solely from the average inclination and is thus relatively independent of the declinations. When the results from the red beds and volcanics are combined the resulting paleolatitude is 20°S , which is indistinguishable from the predicted location of Newfoundland on the North American margin for this time.

We have plotted the paleolatitude of the Springdale Group in Fig. 6 for comparison with predicted paleolatitudes determined from North America, Avalon, and the intervening CMB terranes. All paleolatitudes are calculated for an average Springdale location (49.3°N , 303.7°E). Those from Avalon (the large dots) include values from southern Britain [24,25] that have been calculated from the poles after they have been rotated into North American coordinates so as to close the Atlantic Ocean.

One can see a wide separation between North America and Avalon during the Early Ordovician. This separation gradually diminishes until there is no distinguishable paleolatitudinal differ-

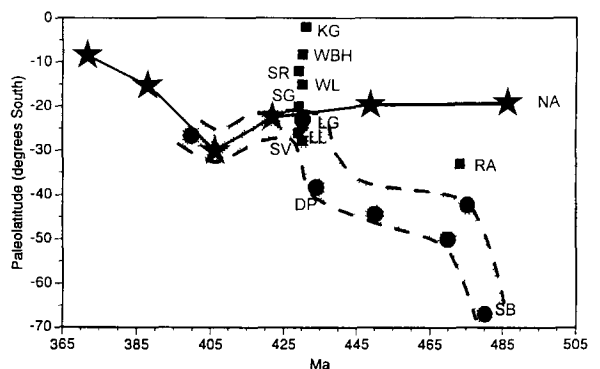


Fig 6 Predicted paleolatitudes calculated for Springdale, Newfoundland (49°N , 303°E) Included are those from central Newfoundland (■), North America (★) and Avalon (●) The Avalon path is highlighted by the stippled area SG = Springdale Group, SR = Springdale red beds, SV = Springdale volcanics, LG = Lawrenceton Formation of Gales et al [5], LL = Lawrenceton Formation of Lapointe [3], KG = King George IV Lake area of Buchan and Hodych [2], WL = Wigwam Formation of Lapointe [3], WBH = Wigwam Formation of Buchan and Hodych [4], RA = Roberts Arm Group of Van der Voo et al [8], NA = North America [22], SB = eastern Avalon (southern Britain) [24,25], DP = western Avalon (Dunn Point volcanics, Nova Scotia) [34].

TABLE 2

Silurian southern hemisphere paleolatitudes from central Newfoundland calculated for the Springdale locality

Unit	Red beds	Mafic volcanics
King George IV Lake	14 ± 9° [26] 2 ± 6° [2]	
Springdale Group	10 + 14 / - 12°	26 + 27 / - 17°
Botwood Group	8 ± 6° [4] 15 ± 9° [3]	24 ± 6° [5] 28 ± 8° [3]

ence in the Silurian. The Roberts Arm Group paleolatitude [8] is shown for comparison because this Ordovician terrane underlies the Springdale Group. The wide spread in paleolatitudes determined from Early Silurian units is clearly apparent. These paleolatitudes range from 2°N for the King George IV Lake area [2] to 27°S for the Lawrenceton volcanics of the Botwood Group [3].

Table 2 lists the paleolatitudes of all Silurian studies of central Newfoundland rocks. The Wigwam Formation of the Botwood Group, studied by Lapointe [3] and Buchan and Hodych [4], consists mainly of red beds. In the former study the Wigwam red beds have shallower inclinations than the Lawrenceton volcanics of the same group (28° vs. 47°). In the King George IV Lake area the mean of three sites of volcanics is 10° steeper than the mean of seven sites of red beds [2]. A more recent study of red beds from the King George IV Lake area yields a low paleolatitude similar to that obtained from the Springdale red beds (14 ± 9°) [26]. Finally, our mean inclination of the Springdale volcanics ($n = 5$ sites) is 25° steeper than that of the Springdale red bed sites ($n = 6$ sites). Hodych and Buchan have reported, in abstract form, similar shallow inclinations from the red beds of the Springdale Group [27]. The only exception to this general pattern is a shallower inclination (by 5°) for three felsic flows included in the Wigwam Formation study of Buchan and Hodych [4]. Although the pattern of shallower inclinations in red beds is fairly consistent, it should be noted that the α_{95} 's are generally large, so that many of these differences in inclination are within the statistical uncertainty. Nevertheless, it remains an important issue because the low red bed paleolatitudes relative to North America have geologic implications which may include thousands of kilometers of south-

ward (dextral) strike-slip motion, which seems geologically implausible.

We will now discuss two categories of possible explanations for the discrepancies in inclinations and their associated paleolatitudes. The first category includes possible problems with the individual studies. The shallower inclinations associated with the red beds may indicate shallowing of an initial DRM, p-DRM or rapidly acquired CRM ($< 10^3$ yrs) due to compaction, as suggested for the Wigwam Formation red beds by Van der Pluijm et al. [28]. Compaction and the associated shallowing of inclination are most likely to occur in fine-grained sediments that are not cemented before significant burial. During paleomagnetic sampling it is common to preferentially sample such fine-grained red beds. In the case of the Springdale red beds the fairly straightforward demagnetization behavior may indicate that NRM acquisition was indeed fairly rapid [29]. Such an early NRM would be susceptible to inclination shallowing upon subsequent burial. In contrast, compaction is negligible in the volcanics of the Lawrenceton Formation studied by Gales et al. [5] and Lapointe [3]. Problems in these studies of volcanic rocks may include the following:

- (1) Paleosecular variation may not be adequately averaged. For mafic volcanics, it has been shown that the actual number of distinct VGPs may be significantly less than the number of flows [30], thus requiring significantly more sites than flows to adequately average paleosecular variation. The very large differences in inclination ($> 30^\circ$), however, cannot be accounted for by this explanation [31].
- (2) The structural corrections from volcanic flows are not as precise as those obtained from sedimentary units (red beds). Imprecise structural correction may account for greater scatter between sites and studies, but not for the consistently higher inclinations exhibited by the volcanics.
- (3) The red beds may have acquired their characteristic magnetization after that acquired by the volcanics. The age of red bed magnetization has been shown to be pre-folding but not necessarily to pre-date the formation of the conglomerates. The conglomerate tests in all studies sampled volcanic clasts and thus limit the timing of magnetization acquisition in the

volcanic units and not the red beds themselves. Thus, any possible remagnetization of the red beds is constrained to be older than the folding, which is pre-Late or Late Silurian in age [32]. Elliot et al. [32] have used U-Pb geochronology to constrain three generations of folds in north-central Newfoundland to be no younger than the cross-cutting tonalite dike swarms (408 ± 2 Ma).

The second category of explanations includes actual latitudinal offset of these terranes relative to North America. If one regards the Lawrence-ton volcanic results as questionable, one may postulate a large dextral offset between the Springdale Group and the King George IV Lake area. A fundamental terrane boundary between the Humber Zone and the CMB must then be postulated between the Springdale Caldera and the King George IV Lake area. Yet the differences seen between red beds and volcanics from the same group (e.g., Springdale and Botwood Groups) would require that the displacement occurred after the volcanics were deposited.

We prefer the explanation of inclination shallowing, because the latitudinal offset hypothesis is geologically less plausible [11]. A mean compaction of the Springdale red beds of 57% would account for the 23° difference in mean inclination between the two rock types. This degree of compaction is comparable with the nearly 50% compaction determined from thin section observations of a highly compacted lithic sandstone [33]. Inclination shallowing in the red beds seems to explain most of the inclination discrepancies between Silurian red bed and volcanic units, and avoids the need for large-scale Silurian displacements in central Newfoundland. This conclusion is also supported by the aforementioned Silurian overlap assemblage [10] in that it would link the CMB and the Humber Zone.

In conclusion, the inclinations derived from the Springdale Group red beds and other Silurian red beds from central Newfoundland should not be used to determine paleolatitudes due to their anomalously shallow inclinations. The nature and occurrence of these inclination discrepancies, however, requires further study within central Newfoundland. The mean paleolatitude derived from the Springdale Group volcanics is 30° , which is indistinguishable from the predicted North

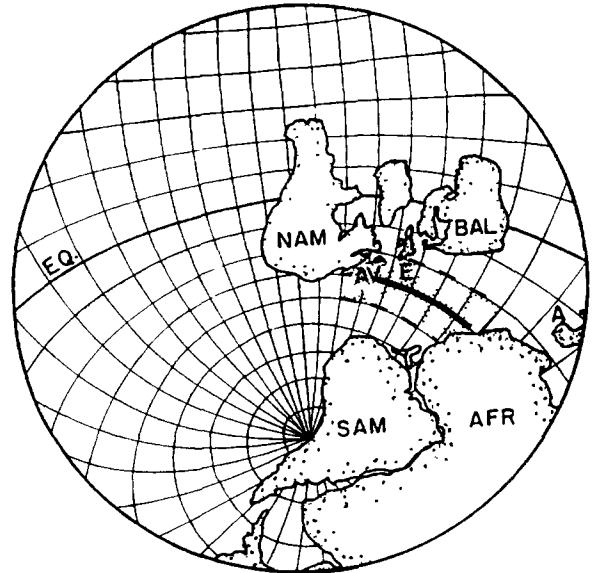


Fig. 7 Silurian paleogeographic reconstruction of the Iapetus Ocean [35]. *A* = Armorica; *AFR* = Africa; *AV* = Avalon; *BAL* = Baltica; *E* = southern Britain; *NAM* = North America; *SAM* = South America. The paleolatitude (30°S) of the Springdale Group is represented by the stippled area, which includes a minimum error estimate of $\pm 10^\circ$. This paleolatitude is indistinguishable from the Laurentian margin and suggests that the western Central Mobile Belt was accreted to Laurentia by the beginning of the Wenlock. The calculated minimum north-south closure velocity for the interval 470–430 Ma is 3.5–4.0 cm/yr, as calculated from the paleolatitudes of the Roberts Arm Group [8] and this study.

American paleolatitude for the Springdale location (Fig. 7). These results imply that the western CMB was accreted to Laurentia by the Middle Silurian and that any post-Silurian latitudinal displacement is beyond the limits of paleomagnetic detection.

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